# BATHYMETRIC MAPPING USING REMOTE SENSING

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by
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FEBRUARY, 1987

### CERTIFICATE

This is to certify that present work entitled,
"BATHYMETRIC MAPPING USING REMOTE SENSING" has been carried
out by Mr. Udaya Singh under my supervision and is not
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February, 1987

(Dr. K.K. Rampal)

Professor

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### ABSTRACT

Bathymetric charting and coastal mapping using MSS data are two parts that constitute this thesis work. The Landsat products such as Computer Compatible Tape (CCT), positive films paper prints were obtained from the National Remote Sensing Agency (NASA), Hyderabad. The area chosen for the study is Andhra Pradesh coastal region near Machlipatnum. The area is covered by one Landsat - 4 scene (Number: 142: 049). The toposheets of this area having the contours of water depth of 9 m and 18 m, were obtained from Survey of India, Dehradun.

A relation is developed between reflectance value (Band and water depth. This relation has been used in preducting the water depth from reflectance value for other areas. Investigat have shown a strong correlation between Band 1 reflectance value (MSS-Landsat 4) and shallow water depths of the sea. The present technique enables us to get an up-to-date hydrographic informat which has been of utmost importance to navigators and coastal engineers. The bathymetric charting by Remote Sensing is not only time saving but also time effective.

### CHAPTER 1

### INTRODUCTION

# 1.1 GENERAL

The last few years have seen tremendous growth in the field of remote sensing. October 1957 marked the beginning of satellite remote sensing with the launch of SPUTNIK by U.S.S.R. For the first time in the history, a man-made object was circling the earth, covering different regions and countrie of the world from its vantage point in space. This led to rap development of satellite borne sensor for reconnaissance purpor Thus remote sensing has advanced from visual analysis of aerial photographs to automatic computer analysis of digital multibal imageries. Today, one has nearly lost count of satellites that spy the earth every moment for military and civilian purposes.

The civilian remote sensing program was essentially an U.S.A. and U.S.S.R. effort in the seventies. U.S.A. made the data available from its satellites to many countries around the world at extremely reasonable costs. As a consequence the use of such data has become common around the world. With the development of high resolution satellites like FRENCH SPOT system, remote sensing techniques are being reliably applied in large number of fields. A number of studies like crop pattern determination, crop estimation, landuse, surface water distributiver course monitoring, forestry planning, geological mapping, are being taken up.

Here in India, we have various facilities like data products laboratory at SPACE APPLICATIONS CENTRE (SAC), full-fledged acquisition and processing facility at NRSA, HYDERABAD for undertaking various projects for diverse applications. The launching of INDIAN REMOTE SENSING SATELLITE (IRS-1), which is scheduled by the end of 1937 will give the indigenous efforts a further boost in the field of remote sensing.

# 1.2 MULTISPECTRAL SCANNER (MSS)

The MSS oscillating mirror scans the information of the object in cross track direction and it travels along the track as shown in Fig. 1.1. It covers an area of 185 km x 185 km in one scene. The information is obtained pixel by pixel and scanned line by line. The total area of 185 km x 185 km covere in one scene is divided into 2340 scan lines and each scan line is divided into 3240 pixels as shown in Fig. 1.2, but there may be small variation in the above numbers. For the data obtained for the present investigation the number of scan lines were 2400 and number of pixels as 3238. Pixel is the minimum unit that can be distinguished from Landsat data. For MSS, the pixe has the dimensions of 57 m x 79 m.

The MSS sensor scans the data in four different bands of the optical and near IR spectrum simultaneously. The bands are numbered as 4, 5, 6 and 7 in Landsats 1, 2 and 3 and as 1, 2, 3 and 4 in Landsats 4 and 5. Each one of the band is useful for

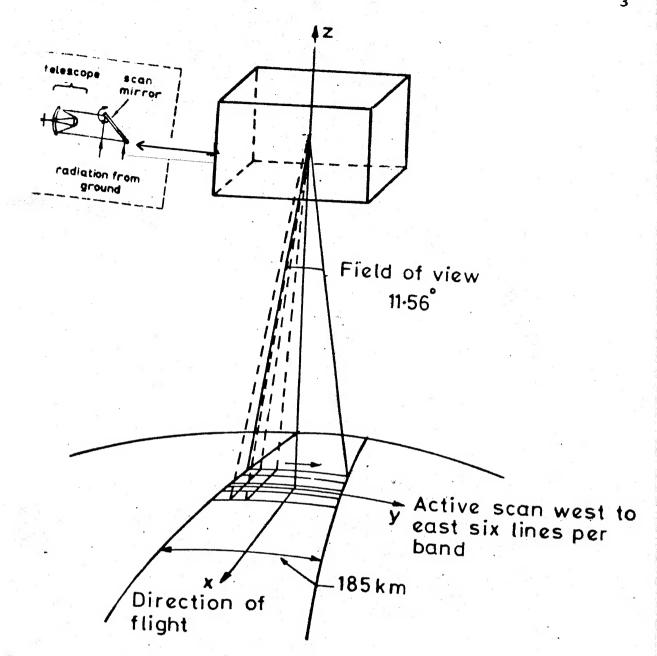


Fig. 1.1 Multispectral scanner

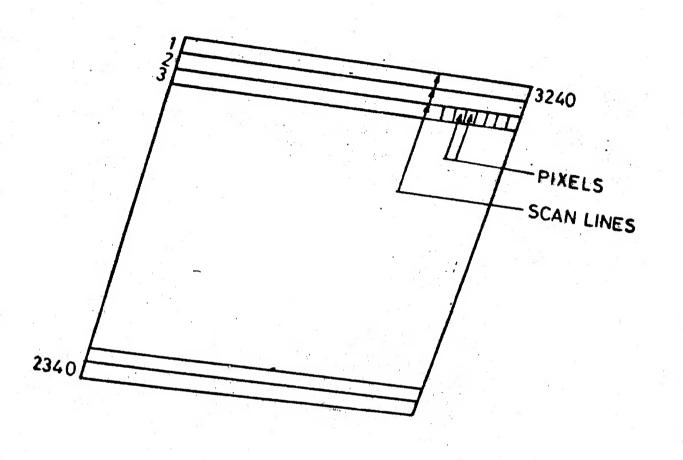


FIG 1.2 SCAN LINES AND PIXELS IN A SCENE

one particular field of study. The various uses of each band are given in Table 1.1.

# 1.3 THEMATIC MAPPER (TM)

It is a special scanner which has better resolution than MSS and operates in 7 channels. Various bands are numbered as 1 to 7 and each one has application in a specific field of study. All the channels have a spectral resolution of 30 m against the 79 m in MSS, except for thermal band i.e. band 7. The various bands and their uses are given in Table 1.1

### 1.4 APPLICATION OF REMOTE SENSING

Remote Sensing has proved its worth in various applications. Indian scientists at various well equipped research organisations and frontline academic institutes have applied the satellite products for diverse applications like geological, agricultural, landuse with diverse techniques like computer aided, visual and image processing.

Snow melt run-off studies (Ramamoorthi, 1983), ground water table prediction (Rampal, 1984), soil mapping (Karale, 1983), forest survey and management (Madhavan Unni, 1983) land evaluation and classification for agriculture (Murthy,

# TABLE 1,1 : SPECTRAL BANDS AND SIGNIFICANCE

Sand to the Sand			
SENSO R	SPECTRAL BANDS	RESOLUTION	APPLICATION
LANDSAT-MSS			
BAND - 1	0.5 - 0.6 um	80 8	Qualitative discrimination of depth and turbidity of standing water bodies.
BAND - 2	0.6 - 0.7 um	80 E	Delineation of topographic and cultural features.
BAND - 3	0.7 - 0.8 um	m 08	Shows tonal contrasts for various landause categories.
BAND - 4	0.8 - 1.1 um	80 m	Land-water discrimination.
LANDSAT THEMATIC MAPPER	TC MAPPER		
BAND - 1	0.45 - 0.52 um	30 m	Increased penetration into water bodies, soil/vegetation and deciduous/coniferou flora discrimination.
BAND - 2	0.52 - 0.60 um	30 m	Vegetation, vigor assessment
BAND - 3	0.63 ~ 0.69 um	30 m	Chlorophyl absorption band for vegeta- tion discrimination, for contrast between vegetation and non-vegetation features.
BAND - 4	0.76 - 0.90 um	30 ш	Biomass content and delineating water bodies.
BAND - 5	1.55 - 1.75 um	30 ш	Vegetation/soil moisture content and snow/cloud differentiation.
BAND - 6	2.08 - 2.35 um	30 m	Discriminating rock types and hydrothermal anomalies.
BAND - 7	10.40 - 12.50 um	120 m	Thermal IR band for vegetation stress, soil moisture and thermal mapping.

Venkataratnam and Saxena, 1983) are few to quote among Indian works.

### 1.5 CRITERIA FOR THE SELECTION OF BAND

It is very important to select a suitable band for a given field of study. The criteria for the selection are as follows.

The imageries which are available in Band 1, 2, 3 and 4 differ in appearance which depend upon the varying reflectance of the different features of the earth's surface. Maximum transmission of light in clear sea water is in the blue spectral region. Water acts as an optical filter, progressively absorbing radiant energy at longer wave lengths until almost complete absorption occurs in the near infrared region. It is shown in Fig. 1.3.

Shallow clear water or water containing suspended materials returns considerable amount of light in less attenuated Band 1 which provides valuable information regarding the depth or nature of foreshore and bottom, concentration of

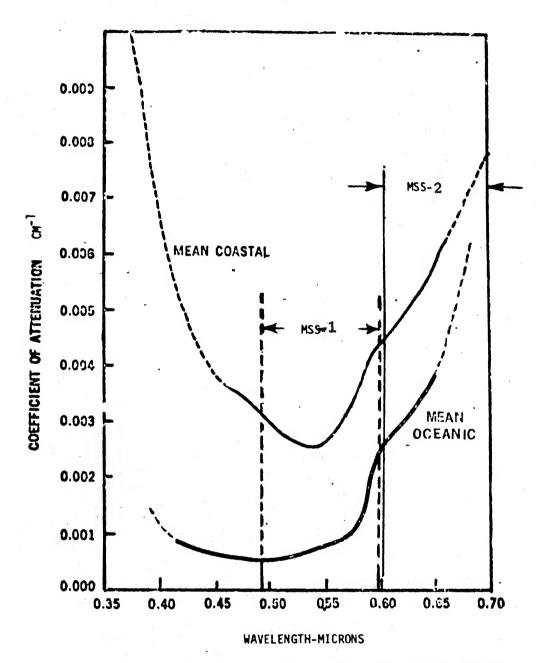


Fig. 1.3: Light Attenuation in Ocean Water in Band - 1 and Band - 2.

suspended material, location of submerged objects.

The attenuation coefficient is higher in band 2, 3 and 4 and very little light is directed towards satellite by water bodies which appears very dark as compared to land surfa Therefore, long wavelength bands particularly band 4 is very useful in distinguishing land from water and are helpful in delineation of coast line, islands, sandspits etc.

Hence for the present work, band 1 was selected for determining water depth and band 4 was selected for mapping th land and water.

### 1.6 NEED OF REMOTE SENSING FOR HYDROGRAPHIC SURVEYING

A study of the hydrographic surveying reveals that except for the extremely small area (5%) surveyed, the remaining areas have been inadequately surveyed, which constitute enormous surveying task. The 1982 Expert Group Report on Hydrographic Perspective Plan has assessed that it would take a span of 10 years by employing 16 surveying vessel and over 4500 personnels to complete navigational surveys and present strength of survey vessels and manpower is deficient. Apart from this, the coastal areas and mouths of rivers on east coast of India like Hugli, Mahanadi, Godavari, Krishna are changing frequently due to floods and storms. The position, extent and depths of many offshore islands, banks, shoals etc.

which are not connected with main land by triangulation are required to be determined. Some ships require navigable area which are more than 22 m deep. An up-to-date hydrographic information has always been of utmost importance to navigators and coastal engineers.

The hydrographers have traditionally relied on sounding machine, echo sounding and aerial photography to maintain and update their nautical charts. Though these inventory technique are quite effective and recent but are slow, cumbersome and costly. At the same time catastrophic events like floods, stor can change the hydrography quickly in delta regions, sandy shore etc. Therefore, the hydrographer requires a method which can give him quick results on bathymetry, coastline, shoals etc. which he is unable to get by conventional method, without losing much time.

In the continued attempt to introduce advance technology fitting to the navigator's requirement attempts have been made to develop the new method of collecting offshore hydrography through remote sensing techniques. This method has greatly redute cost and improved the quality of work. This has great potential to collect large amount of hydrographic data very rapidly over shallow coastal water.

In this present project, an attempt has been made to correlate the reflectance of water bodies obtained by sensor and water depth for shallow water. We have prepared the map of coastal area distinguishing land, clear water and water having suspended materials. Our area of study in Andhra Pradesh coastal area between Madras and Kakinada. Details of the study are explained in the following chapters.

### 1.7 OBJECTIVE OF THE STUDY

The objective of the study is:-

- (1) to correlate the reflectance (MSS Band 1) with shallow water depth
- (2) to map the coastal area distinguishing land, water having suspended materials and clear water.

To achieve the first objective of the work, it is necessary to get water depth contours of that area, to calcul the latitudes and longitudes of the points on the known water depth contours, to convert these coordinates to line numbers and pixel numbers and to read the corresponding reflectance values from CCT obtained from NRSA in MSS Band 1. The procedure for this work is given in chapter 3.

For the second objective an area, which covers the land, water containing suspended materials and clear water, was chosen on imagery and reflectance values of all the pixel falling over that area were read from CCT. Maps were prepare by the program which is given in Appendix 6.

### CHAPTER 2

### SCOPE OF THE WORK

### 2.1 EARLIER STUDIES

Not much literature is available on the application of Remote Sensing in the field of measurement of water depth. Bathymetry is very old field for navigators and they are using so many traditional methods of water depth measurement like sounding machine, echosounding, aerial photography etc. Thoughthese techniques are quite effective but are slow.

The work of Ross<sup>(2)</sup> and Polcyn and Rollin<sup>(8)</sup> have shown that there is a correlation between MSS band 1 imagery a shallow water depth. Correlation was found between MSS - 1 imagery and depth in number of areas corresponding to water depth of less than 2 m, 5 to 10 m. But the depth predicted from imagery was approximate. Under favourable condition Landsat imagery can detect sub-surface detail as small as 200 in diameter. Water depth over a large area of world coastline and shallow sea can be assessed by Landsat imagery and image processing technique which would be valuable aid to chart revision by directing the hydrographer to area requiring close examination.

The wave refraction method was applied by Polycn in 1969. The underlying principle in this case is that when a wave moves through a liquid the characteristics of the medium which may influence its motion are:-

- (a) The depth and other boundary conditions.
- (b) Gravitation, since the changed profile or contour of the surface involves work against gravity.
- (c) Surface tension, because the pressure under a curved surface is different from that beneath a flat surface.
- (d) Viscosity, which is the dissipative energy agent.

The velocity of propagation for a surface wave in a liquid of density e is given by

$$v^2 = \frac{2\pi S}{\rho \lambda} + \frac{g\lambda}{2\pi}$$
 (2.1)

where S = surface tension

g = acceleration due to gravity

 $\lambda = \text{wave length}$ 

For sufficiently large value of  $\lambda$  i.e., for long waves first term is negligible compared to second term of the above expression. Thus the velocity of 'Gravity Waves' is given by

$$\mathbf{v^2} = \frac{g \lambda}{2 \pi}$$

These gravity waves are observed over deep part of the water body. We know that

$$v = f\lambda$$

where, f = frequency of wave.

Since time period of a wave is the inverse of the frequency i.e. time period  $(p) = \frac{1}{f}$ 

Finally we get the relation between velocity and time period as

$$v = \frac{\gamma}{p} = 5.09 p$$

where  $g = 32 \text{ ft./sec}^2$ 

v being measured in ft./sec.

The velocity in the shallow water is given by the expression

$$v^2 = gd (2.2)$$

where v = velocity of the wave

g = gracitational acceleration

d = depth of water.

As the wave moves from deeper portion to the shallow part, the time period of the wave 'p' remains constant. This property is utilized to determine the depth. If  $\mathbf{v}_{o}$  be velocity for deep water, then we have

$$v_0^2 = \frac{g \lambda_0}{2 \pi} = (f_0 \lambda_0)^2 = \frac{\lambda_0^2}{p^2}$$

where f = frequency of wave in deep water

 $\gamma_0$  = wave length in deep water

 $p = time period of wave = 1/f_0$ 

Therefore, 
$$p^2 = \frac{2\pi \lambda_0}{q}$$
 (2.3)

For shallow water

$$v^{2} = gd = \frac{\lambda^{2}}{p^{2}}$$

$$p^{2} = \frac{\lambda^{2}}{gd}$$
(2.4)

Therefore,

Now from equations 2.3 and 2.4, we get

$$\frac{\lambda}{\lambda_0} = \frac{2\pi d}{\lambda} \tag{2.5}$$

and also

$$\frac{\mathbf{v}}{\mathbf{v}} = \frac{\lambda}{\lambda_0}$$

The ratio  $\frac{\lambda}{\lambda_0}$  is thus functionally related to the ratio  $\frac{d}{\lambda}$ 

A measurement of  $\gamma_o$  &  $\gamma$  will then allow 'd' to be determined. This can be done by successive picture at short intervals to obtain the values of  $\mathbf{v}_o$  &  $\mathbf{v}$ . Measurement of  $\gamma_o$  can be made directly from an aerial photograph. This method can, therefore, be applied easily where seashore or the bed of the water body slopes gently.

Polcyn & Lezenga (3) demonstrated the use of satellite data for mapping of shallow water features for the purpose of upgrading the world navigation charts. A mathematical model for depth measurement using the ratio of reflectance in Band 1 and Band 2 had been successfully developed. Satellite data also provide geographical evidence for verifying existence or non-existence of doubtful shoal features appearing in the world charts and considered to be hazardous to shipping.

The technique employs the relation shown by equation

$$Z = \frac{1}{f(\Theta, \emptyset) (K_1 - K_2)} \ln \frac{\alpha_1 V_1 H_1 \ell_1}{\alpha_2 V_2 H_2 \ell_2}$$
 (2.6)

Where Z = water depth.

 $K_1, K_2$  = attenuation coefficients of water in two different wave lengths.

 $P_1, P_2$  = reflectances for bottom material in two different bands.  $\alpha_1, \alpha_2$  = constants of instrument which are known.  $H_1, H_2 = incoming solar radiations.$ 

 $V_1, V_2$  = analog signals observed in the multispectral scanning.

 $\Theta$  = observation angle.

 $\emptyset$  = solar zenith angle.

Here  $V_1$  and  $V_2$  are obtained from Band 1 and Band 2 of CCT. The incoming solar radiation  $(H_1,\ H_2)$  is available from standard references.  $\mathcal{C}_1$  and  $\mathcal{C}_2$  are reflectances for bottom materials, which are known.

Photogrammetric bathymetry is one of the traditional techniques of remote bathymetry. Colour aerial photography with its remarkable clear water penetration characteristics and dramatic presentation of submerged detail is basic tool in this method for mapping seabed in water of moderate depth. Aerial photogrammetry of bathymetry requires several minor departures from the method normally used where bundle of light rays passes only through atmosphere. The effect of refraction at the water air interface must be taken into account in the aerotriangulation for imaged point. The solution of this problem requires a mathematical model for two media refraction.

To get correct and entirely acceptable for photogrammetric trangulation, the mathematical modelling must be based upon the actual under water position of the point rather than its refrected

or apparent position. The basic mathematical model for correcting image coordinates of under water points is illustrated in Fig.2.1

$$\Delta d = d_{a}h (1 - 1/a)/(H - h)$$
 (2.7)

where  $\triangle d$  = the correction in meters for image point. Its sign is always negative (correction towards the photocentre).

d = radius of image point in the form  $(x^2 + y^2)^{\frac{1}{2}}$ 

h = depth in meters of the underwater point at the time of photography.

H = flying height in meter.

a = ratio of tangents of angles of refrection(r) and incidence (i).

i.e. 
$$a = \frac{\tan r}{\tan i} = [\mu^2 + (\mu^2 - 1) \tan^2 r]^{\frac{1}{2}}$$

where,  $tan r = \frac{d}{f}$ , for vertical photograph.

 $\mu$  = index of refrection for ray passing from water to air f = camera focal length in meters.

If 
$$\mu = 1.34$$

$$a = [1.7956 + (0.7956) d^2/f^2]^{\frac{1}{2}}$$
 (2.8)

The relation between refraction corrected coordinate value

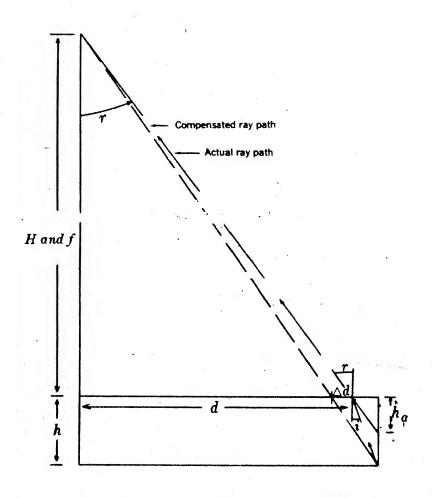


Fig.2.1: Image Displacement due Water Surface Refraction

(x', y') and the photo image coordinate (x, y) is shown below

$$x' = x(1 - \Delta d/d)$$
  
 $y' = y(1 + \Delta d/d)$  (2.9)

The change in the value of equation (2.8) and its effect on  $\triangle$ d caused by 1° or 2° tilt is insignificant in aerial photograph if H is greater than 100 h. Therefore, the refrection compensated coordinates of all the vertical control point can be determined. To be theoretically correct, the unknown depths (h) of the test point should be solved as additional unknown in the aerotriangulation process using equation (2.7) through the iterative process.

Moore (1947) developed a transparency method for water depth from aerial photography alone. This was one of the first remote sensing method. Basically the method consists of determining the optical extinction coefficient of the water column from the brightness measured at identical points in two specially filtered (red and green) photographs. The spot depths are determined by means of a ratio calculator which compares the brightness in two band to know water colour-transmission. Results with ± 10% accuracy to a depth of about 6 m are achieved over homogeneous san bottoms. This technique is limited to calm sea, relatively clear water, bright skies, without cloud patterns and sun angle between 30° and 55°.

### 2.2 AREA UNDER STUDY

From the work of Ross<sup>(2)</sup>, Polcyn and Lyzen<sup>(3)</sup>, Jha and Satya Prakash<sup>(5)</sup> and others it is clear that MSS band 1 reflectance values could be used to predict shallow water depth. Satisfying above condition, coastal area of Andhra Pradesh was selected. The area is along the east cost of the Andhra Pradesh region between Madras and Kakinada.

The area is covered by one Landsat - 4 scene. For reference the index map of Landsat - 4 coverage is shown in Fig.2. The index map is supplied by NRSA. Any scene is located by Path number and Row number. The area selected is located by Path number - Row number: 152 - 049. The area inclosed by the longitudes and latitudes is shown in Fig. 2.3.

# 2.3 DATA COLLECTED FOR THE WORK

The data collected for the work are as follows:

- (1) Landsat data.
- (2) Topo sheets with water depth contours.

The Landsat data in the form of CCT containing reflectance values for band 1, 2, 3, 4, positive films(1:1 M) and paper print (1:1/2 M) were obtained from NRSA, Hyderabad, While selecting

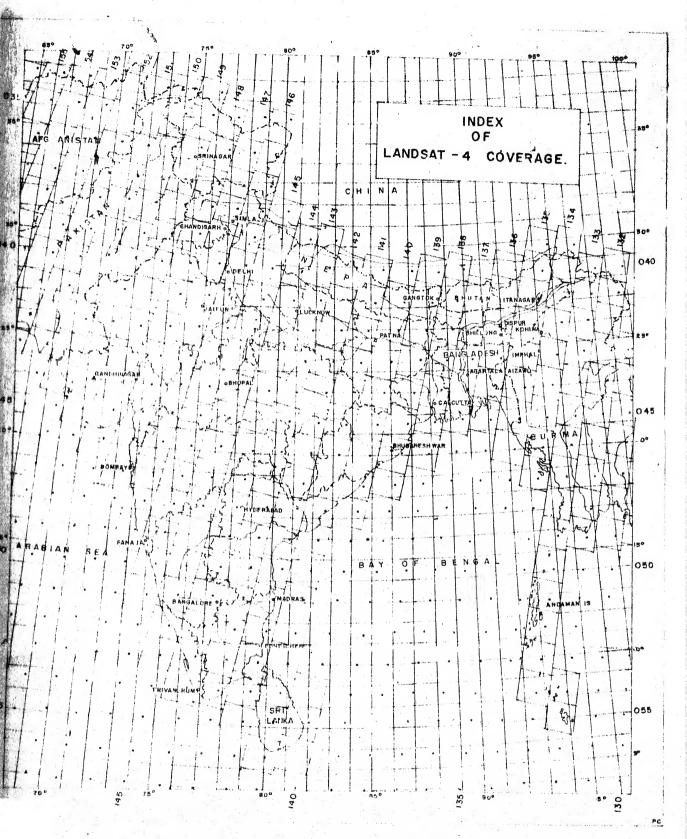


Fig. 2.2: Index Map of Landsat 4 Coverage (Supplied by NRSA)

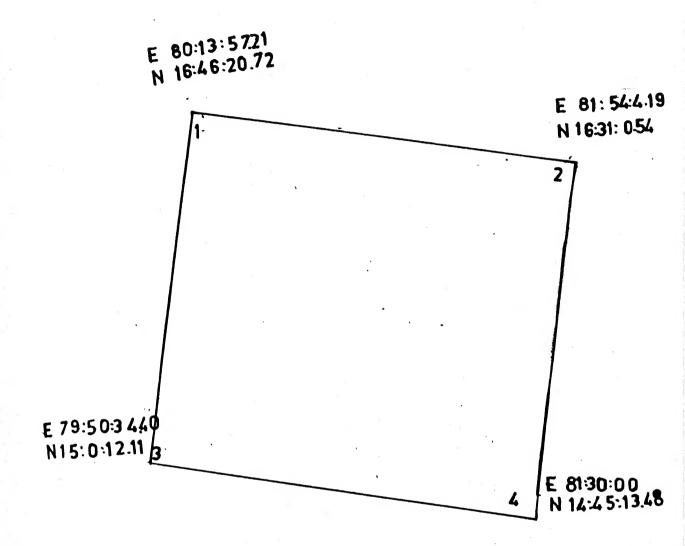


FIG 2.3 AREA UNDER STUDY

is exercised so that there is no cloud cover. The date of pass for present data is for 28th September 1983.

The toposheets of the area were obtained from SURVEY OF INDIA. The area is covered by two toposheets of scale 1:250,000 (1:1/4 M). With reference to index map of Survey of India, the two maps are designated as 66 A & E and 65 H. To obtain the information about availability of map of required scale, costs, whether restricted or not one may write to MAP SCALES OFFICE, HUDA COMPLEX, TARNAKA, HYDERABAD. On these two toposheets itself, the two water depth contours of 9 m and 18 m are shown.

### CHAPTER 3

### METHODOLOGY

### 3.1 INTRODUCTION

To obtain the reflectance value of a ground point, it is necessary to convert the longitude and latitude of that ground point to corresponding line number and pixel number. For this it is necessary to determine all the corner coordinates of the imagery very precisely. The corner points coordinate are as shown in Fig. 3.1.

# 3.2 CONVERSION OF GEOGRAPHICAL COORDINATE TO CONICAL ORTHOMORPHIC COORDINATE

Projection of geographical coordinates of Landsat imagery which is on curved surface is to be converted to a flat surface whose X and Y coordinates can be computed. This map projection used for this purpose is the conical orthomorphic projection which is used country like India to minimize projection distortion

The orbit of Landsat makes an angle about 9° with the North direction. Hence the edges of image are not parallel

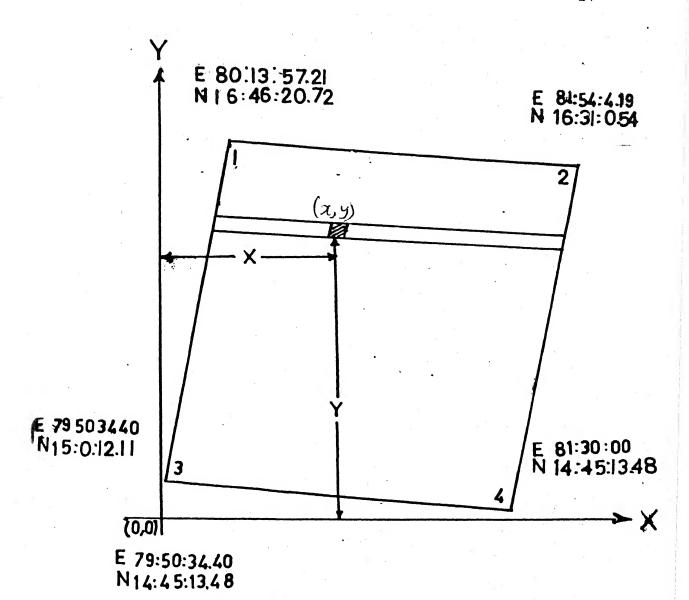


FIG 3.1 CORNER COORDINATES OF IMAGE AND CONVESION TO LINE NO. AND PIXEL NO.

to the longitude lines. The geographical coordinates of the corner points are converted to conical orthomorphic coordinates with reference to chosen coordinate axis system parallel to longitude and latitude. The origin is chosen to be a point just outside the imagery preferably just below it and slightly towards the left so that no coordinates (X, Y) become negative as shown in Fig. 3.1.

For the conversion of geographical coordinates to conical orthomorphic coordinates, latitudes and longitudes of the four corners of imagery should be precisely known. These value can be obtained from the records published by the Eros Data Centre, Sious Falls, South Dakota (The center supplies Landsat-imageries and Tapes on a world-wide basis). Fig. 3.2 represents the four corners of the imagery. The origin 0 is chosen to be point outside the imagery as reference point. Let the origin 0 has geographical coordinates ( $\emptyset_0$ ,  $L_0$ ) where  $\emptyset_0$  is latitude and  $L_0$  is longitude. Again let ( $\emptyset_p$ ,  $L_p$ ) be geographical coordinates of any point on the imagery. The conical orthomorphic coordinate can be given as:-

$$X = (P - m') \sin \gamma$$

$$Y = m' + X \tan \left(\frac{\gamma}{2}\right)$$

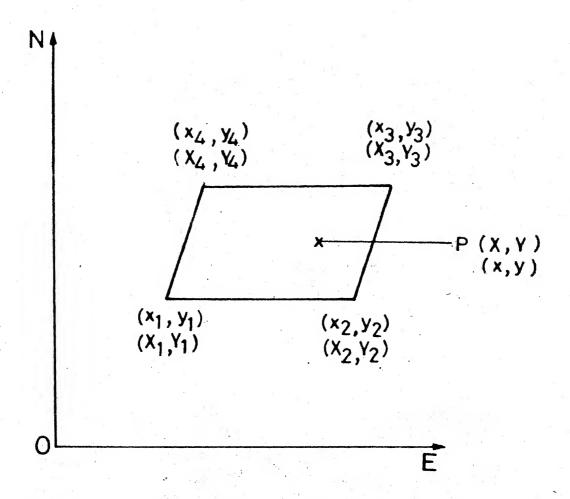


Fig. 3.2 Transformation of geographical co-ordinates to conical orthomorphic co-ordinates

where, 
$$\gamma = \Delta L \sin(\frac{\emptyset_0}{2})$$

$$\Delta L = L_p - L_o$$

$$m' = m + \frac{m^3}{6R_0N_0} + \frac{m^4 \tan \frac{\emptyset}{0}}{24 R_0N_0^2} + \frac{m^5 (5 + 3 \tan^2 \frac{\emptyset}{0})}{120 R_0N_0^3}$$

 $R_{o}$  = Radius of curvature of the earth (Meridianal)

$$= \frac{a(1 - e^2)}{(1 - e^2 \sin^2 \phi_0)^{3/2}}$$

a = major axis of the reference ellipsoid

= 6377277.6 m (For Everest Ellipsoid adopted for India)

b = minor axis of the reference ellipsoid.

= 6356075.0 m (For Everest Ellipsoid adopted for India)

$$e^2 = eccentricity = 1 - b^2/a^2$$

For the Everest Ellipsoid,  $e = \frac{1}{300.8}$ 

 $N_{o}$  = normal to the surface at  $\emptyset_{o}$  (Normal section radius)

$$= \frac{a}{(1 - e^2 \sin^2 \varphi_0)^{\frac{1}{2}}}$$

$$m = R_{m}(\emptyset_{p} - \emptyset_{o})$$

 $R_{m}$  = Mean radius of the earth (Normal section radius)  $= \sqrt{N_{0}R_{0}}$ 

A program to calculate the conical orthomorphic coordinate of a point given its longitude and latitude is given in the Appendix - 1.

Here origin of the reference axis was taken just below the third corner point and has its coordinate E 79° 50' 34.40", N 14° 45' 13.48". The corner points 1, 2 and 3 were converted to conical orthomorphic coordinates. The conversion results are shown in Table 3.1.

## 3.3 CALCULATION OF THE CONSTANT OF THE AREA

Transformation from conical orthomorphic coordinate to line number and pixel number can be made by the following equations.

$$X = A_1 x + B_1 y + C_1$$

$$Y = A_2 x + B_2 y + C_2$$

where X = conical orthomorphic coordinate of the point:Longitude
Y = conical orthomorphic coordinate of the point: Latitude

x = Line number

y = Pixel number

Corner Point	Geographical coordinate		Conical orthomorphic coordinate				
POINC	Longitude	Latitude	Longitude (X)	Latitude (Y)			
1.	80,232558	16.772421	41552,5824008	224769.6131003			
2.	81.901163	16.516816	219743.0386210	197273.3419803			
3.	79.842890	15,003363	0,000	27783,9174497			

Now we have six equations and six unknowns  $A_1$ ,  $B_1$ ,  $C_1$ ,  $A_2$ ,  $B_2$  and  $C_2$ . These six simultaneous equations were solved and unknowns were determined.

The calculated values of the constants are:-

$$A_1 = -17.32079300$$
 $B_1 = 55.04802500$ 
 $C_1 = 41514.855000$ 
 $A_2 = -82.111587$ 
 $B_2 = -8.4943687$ 
 $C_2 = 224860.220000$ 

A<sub>1</sub>, B<sub>1</sub>, C<sub>1</sub>, A<sub>2</sub>, B<sub>2</sub> and C<sub>2</sub> are constants whose values can be determined from a minimum of three ground points. The three points chosen are three corners of the imagery whose (x, y) coordinates are already computed from early transformation from their geographical coordinates. The (x, y) coordinate (line number and pixel number) of these three points are also known, since they are corner of the imagery. The set of equations are:

$$X_1 = A_1 X_1 + B_1 Y_1 + C_1$$
 $Y_1 = A_2 X_1 + B_2 Y_1 + C_2$ 
 $X_2 = A_1 X_2 + B_1 Y_2 + C_1$ 
 $Y_2 = A_2 X_2 + B_2 Y_2 + C_2$ 
 $X_3 = A_1 X_3 + B_1 Y_3 + C_1$ 
 $Y_3 = A_2 X_3 + B_2 Y_3 + C_2$ 

Here suffixes 1, 2 and 3 refer to corner points 1, 2 and 3.

# 3.4 DETERMINATION OF LINE NUMBER AND PIXEL NUMBER

Now the line number and pixel number of any point on the imagery can be found by solving the two simultaneous equations

$$X = A_1 x + B_1 y + C_1$$

$$Y = A_2 x + B_2 y + C_2$$

Here  $A_1$ ,  $B_1$ ,  $C_1$ ,  $A_2$ ,  $B_2$ ,  $C_2$ , X and Y are known and X and Y can be calculated. A program for solving these equations thereby determining the line number and pixel number of any point on the imagery is given in Appendix - 2.

By giving geographical coordinate in degree, minute and second, the program calculates the line number and pixel number of any point within that area.

The scale of toposheets is 1:250,000. The latitude and longitude of all the necessary points was determined precisely with toposheet.

# 3.5 READING THE COMPUTER COMPATIBLE TAPE (CCT)

The input data for the present study were taken from CCT. Each pixel is encoded in a byte on the CCT. Each byte is composed of eight binary digits (bits) which are arranged to represent different brightness value as binary number.

The data is stored in CCT in three files; 1st and 2nd files being Tape Directory and Header Record respectively. The tape directory identifies the contents, format of CCT etc. The header record gives the content of data and discribes the format

in which the data are recorded. Third file contains the data in four bands. Third file consists of 9600 records and each record contains 3596 bytes. Each record in third file represents one scan line in one band. Each scan line requires four records to be completely described in all four MSS bands. The mode of data storage is Band-Inter-Leaved (BIL). In this type of data storage, first four records correspond to scan line number 1. One for each band i.e. first record gives the data for line number 1 in band 1, second records gives the data for line number 1 in band 2, third records gives the data for line number 1 in band 3 and fourth record gives the data for line number 1 in band 4. Fifth record to eighth record corresponds to scan line number 2 in all four bands, i.e. fifth record corresponds to band 1, sixth record corresponds to band 2 and so on, in this tape, to get the data of xth line in band 4, one has to read '4x'th record. In each record before and after the data, there are zero fills. The number of zero fills varies from record to record. If there are n initial zeros in the record and if we want the reflectance value of Nth pixel, then we have to read the reflectance value of (N + n) th pixel (Byte).

The CCT used in DEC-10 system has 9 tracks. A frame consists of bits 0 to 7 in which the pixel brightness is recorded and a ninth parity bit. A word is made of 4 bytes or

36 bits. This is shown in Fig. 3.3.

A program for reading CCT is given in Appendix - 3.

This program reads one scan line in one band of the image file. The CCT as obtained could not be read on DEC-10 system. This is because DEC-10 is 36 bits word machine and CCT is formated for 32 bits word machine. If one tries to read the CCT as obtained the DEC-10 system will not read the 5th, 10th, 15th......bytes due to lack of compatibility between the tape format and DEC-10 system. A dummy blank was introduced after 4th, 9th, 14th.....bytes increasing the record size from 3596 bytes to 4500 bytes. When the tape is read on DEC-10 system, the introduced dummy blank bytes will be skipped and data of all 3596 bytes will be printed.

The required record number to be read, is found by multiplying the line number with 4 and adding the band number. All the record numbers are stored in the increasing order and CCT is read record by record. To get the reflectance of a pixel, the number of initial zeros are counted in each of the record and added to pixel number to get the byte number corresponding to that pixel and the reflectance value is read from CCT for that byte.

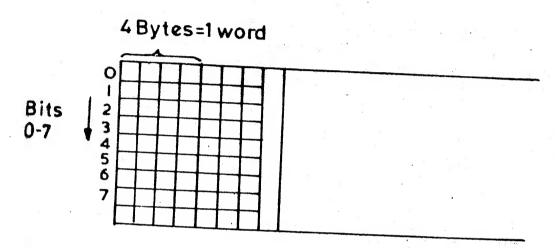


Fig. 3.3 Cross section of a 9 track CCT

### CHAPTER 4

#### NUMERICAL RESULTS

### 4.1 RELATION BETWEEN REFLECTANCE VALUE AND WATER DEPTH

In water area, the light reaches back to MSS (sensor) as reflection from water surface, scattering directly in atmosphere, reflection from bottom (seabed) and reflection or scattering from the object inside the water. Scattering directly from atmosphere and surface reflection is same for same condition. But reflection from bottom is affected by suspended material which can scatter the light. The amount of light returned from the bottom depends upon the depth of water, its attenuation coefficient for the wave length and reflection coefficient of the bottom. The light reaching the bottom is also affected by the slope angle E' which can be calculated, if refractive index of sea water is known. Light path from sun to satellite is shown in Fig. 4.1.

For water depth d, the optical path length in water

$$= d(1 + cosec E')$$
 (4.1)

If  $\mu$  is refractive index of sea water then

$$\mu = \frac{\sin (90^\circ - E)}{\sin r} \tag{4.2}$$

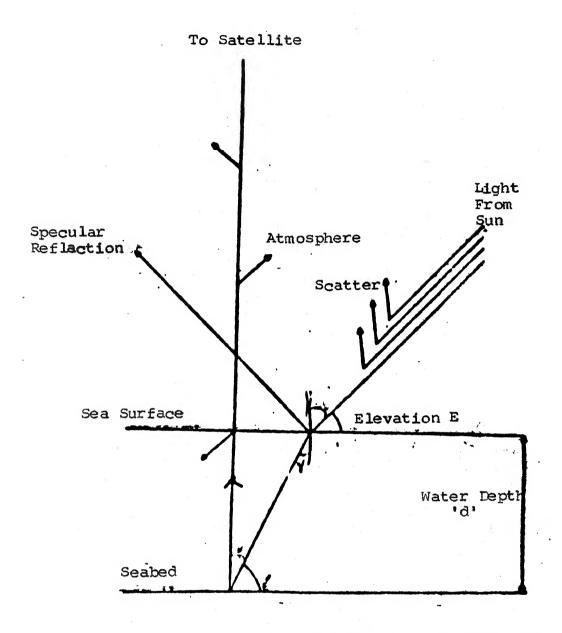


Fig. 4.1 : Light Paths From Sun to Satellite

where, E = sun elevation

$$r = \sin^{-1} \left( \frac{\cos E}{\mu} \right) \tag{4.3}$$

$$E' = 90^{\circ} - r \tag{4.4}$$

Here  $E = 55^{\circ}$  and  $\mu = 1.34$ 

Therefore E' = 64° 39' 23"

The value of sun elevation E, is given in imagery of study area. When light passes through a distance dL in water, then energy is attenuated to an amount dE, which is proportional to the energy E and the distance dL

or 
$$dE = -K_*dL_*E$$
 (4.5)

where K = attenuation coefficient and is dependent on wave length. The negative sign indicates loss of energy. If  $E_{\rm O}$  and  $E_{\rm L}$  be energy of radiation for L = 0 and L = L respectively.

Then,

$$\int_{E_{O}}^{L} \frac{dE}{E} = - \int_{O}^{L} KdL \qquad (4.6)$$

$$E_{L} = E_{O} \cdot e^{-KL} \tag{4.7}$$

i.e.  $E_0$  is the energy of light when it just inters the surface. L is optical path length in water. Here L = AB + BC

or 
$$L = d(1 + cosec E')$$
 (4.8)

Therefore 
$$E_L = E_0 \cdot e^{-Kd} (1 + cosec E')$$
 (4.9)

Energy received by the sensor will be sum of energy directly coming to sensor due to atmospheric scattering and refrected energy coming from water  $E_o$ .  $e^{-Kd(1 + cosec\ E')}$ .

If  $\mathbf{E}_{\mathbf{T}}$  be the total energy received by the sensor then

$$E_{T} = Ed + AE_{o} e^{-Kd(1 + cosec E')}$$
 (4.10)

where  $E_d$  = energy received directly through scattering by sensor and is constant for same conditions.

A = parameter depending upon state of atmosphere.

or 
$$E_T - E_d = AE_o e^{-Kd (1 + cosec E')}$$
 (4.11)

If  $E_x$  and  $E_y$  be the energy (Reflectance value) received by sensor for water depth x and y respectively, then

$$\frac{E_{x} - E_{d}}{E_{y} - E_{d}} = e^{-K (1 + \csc E') (x - y)}$$

$$\frac{\ln(\frac{E_{x} - E_{d}}{E_{y} - E_{d}})}{-K(1 + \csc E')}$$
or
$$x - y = \frac{-K(1 + \csc E')}{-K(1 + \csc E')}$$
(4.12)

If y = 0, then

$$depth x = \frac{\ln(\frac{E_x - E_d}{E_o - E_d})}{-K (1 + cosec E')}$$
(4.14)

Where  $E_0$  is energy (Reflectance value) received by MSS, when water depth  $~\sim~$  O.

In the above equation  $E_d$ ,  $E_o$  and K are unknowns. For the determination of  $E_d$ ,  $E_o$  and K, 30 points were taken on the toposheets which lie on the known depth countours. Latitudes and longitudes of all the points were calculated precisely. Their line number and pixel number and finally the corresponding reflectance value were read from CCT which is shown in Table 4.1. Now we have thirty equations and three unknowns.  $E_d$ ,  $E_o$  and K were calculated by least Square Method.

Solution by least square is given by the equation

$$X = - (A^T PA)^{-1} A^T PL$$

where A = coefficient matrix

= matrix of partial derivative of the function with respect to unknowns.

P = weight matrix.

Here weight matrix is taken as Unit Matrix.

$$L = (L_0 - L_b)$$

POTATS ARE TAKES ON 18m WATER DEPTH CONTOUR

	31.	LATITUDE	LONGITUDE	LINE	PTXEL No.	PEFLECTANCE VALUE
ing amos	1 6	15.49611	80.50000	1624	1037	28
	?.	15.25361	80.25000	1978	662	28
	3.	15.50000	80.50694	1617	1049	30
	A	15.25000	80.23833	1998	645	28
		15.60917	90.39472	1496	791	28
	6.	15.32028	80.33500	1887	799	28
	7	15.17306	80.20694	2106	619	30
	8.	15.13861	80.13917	2164	505	27
	ο,	15.02078	80.17028	2310	612	28
	1 .	16,13333	81.53333	579	2712	28
	11.	15.93333	91.33333	882	2427	28
	12.	15.73333	81.13333	1185	2131	27
	13.	15.63333	80.66667	1542	1336	27
	10.	15,40000	80.45657	1756	1015	31
	15.	15.20000	90.25667	2059	720	2.8

TABLE: 4.1 (....CONTINUED)

## POINTS ARE TAKEN ON 9m WATER DEPTH CONTOUR

	31.	IATITOE	GONGTTUDE	LINE	PIXEL NO.	REFLECTANCE VALUE
ADS WALK	40	15,26667	81.53333	404	2655	37
	?.	10.16657	91.33333	576	2322	37
	₹.	16.00000	81.26667	808	2268	36
	4.	15.86667	<sup>9</sup> 1.13333	1010	2074	35
	4.1 m	15.66667	91.00000	1300	1908	37
	Fin	15.66667	80.73333	1353	1405	36
	"7 .	15.66361	80.77667	1349	1488	36
	Q .	15.75000	80.37833	1315	702	35
	٥.	15.66333	80.91556	1321	1750	36
	10.	15.79667	80.67556	1194	1242	37
	11.	15,47028	80,35722	1686	779	37
	12.	15,51472	80.48778	1602	1006	37
	13.	15.36139	80,25000	1850	622	36
	14.	15.25000	80.10972	2024	403	35
	15.	15.02333	90.06778	2330	418	37

L<sub>o</sub> = Matrix of computed value of x (depth) by assumed unknowns

 $L_b = Observed value of x (depth)$ 

X = Correction Matrix for unknowns.

The expression can be write more simpler as

$$X = - (A^{T}A)^{-1} A^{T}L$$

$$X_a = [X_o + X]$$

where, X = Matrix of adjusted value of unknowns

X<sub>o</sub> = Matrix of assumed value of unknowns.

For the assumed values of unknowns K,  $E_{\rm o}$  and  $E_{\rm d}$ , correction matrix X was determined as explained above and correction matrix was added to assumed values of unknowns. This process is repeated till we get very less variation in the correction matrix.

A program for least square method is given in Appendix - 4 and value of unknowns are as follows:

$$E_d = 23.61$$

$$E_0 = 62.18$$

$$K = 0.057 \text{ m}^{-1}$$

# 4.2 CALCULATION OF WATER DEPTH

Once the unknown  $E_{d}$ ,  $E_{o}$  and K are determined, then it is very easy to calculate water depth at any point by equation 4.14, if reflectance value of that point is given.

## 4.3 DISCUSSION OF RESULTS

There are only two depth contours on the toposheets. For checking the accuracy of the work, 15 points on each of the contour lines were taken. Latitudes and Longitudes of all those points were calculated. Line numbers and pixel numbers of those points were calculated and reflectance values of all those points were read from CCT. Water depths of all those points were calculated by equation 4.13. The calculated depth of all thirty points are shown in Table 4.2. Then the difference of actual water depth taken from toposheet and calculated water depth is calculated for both the set of data 9 m and 18 m water depths separately.

Standard error (Se) and coefficient of correlation (r) are calculated as:

Se = 
$$\sqrt{\frac{\sum (Y_i - Y_{est}^2)}{N}}$$

$$r = \frac{n\sum xy - (\sum x) (\sum y)}{\sqrt{\left[n\sum x^2 - (\sum x)^2\right] \left[n\sum y^2 - (\sum y)^2\right]} }$$

TABLE :4.2

HE "HE GIVEN DEPTH(Y)=18m

est=DEPTH CALCULATED

1.	TANTEROR	Ponditabe	LINE No.	PIXEL,	REFLE- CTANCE	Yest	Y=Yest
nice ages total a	15,50000	80,50694	1617	1047	28	18.1	-0.1
	15.52278	80.50000	1589	1026	28	18.1	-0.1
3 .	15.68861	81.00000	1271	1898	27	20.3	-2.3
	15.51866	80.81056	1402	1507	28	18.1	-0.1
۲, "	15,60306	80.62750	1458	1233	2.8	18.1	-0.1
5.	15.32250	80.36805	1878	861	29	16.4	1,6
7.	15,21139	80.21972	2053	627	27	20.3	-2,3
23 4	16,13333	80.60891	566	2838	28	18.1	-0.1
er.	15.83333	81.26667	1027	2339	29	16.4	1.6
	15,63333	80.60211	1555	1210	29	16.4	1.6
4	15.43333	80.53333	1699	1126	28	18.1	-0.1
.2.	15.33333	80,40211	1857	917	28	18.1	-0.1
3.	15.21139	80,21972	2053	627	27	20.3	-2.3
0.	15,62028	80.77677	1405	1507	28	18.1	-0.1
5.	15.25000	80,23833	1998	646	28	18.1	-0.1
-							

TABLE :4.2(.....CONTINUED)

FIR THE GIVEN DEPTH(Y)=9m

Yest=DEPTH CALCULATED

1.	LATTITIES	LONGITUDE	LINE.	PIXEL	REFLE- CTANCE	Yest	Y-Yest
TOTAL TOTAL DESIGN IN	15.75000	80,.71500	1248	1336	37	8.8	0.2
2.	15.64639	80.85861	1335	1650	37	8.8	0.2
١.	15.73667	80.72972	1262	1369	36	9.5	-0.5
1.	15.68861	80.34472	1402	644	35	10.2	-1.2
5.	15.19528	80.34583	1655	747	37	8.8	0.2
6.	15.57914	80,54889	1505	1094	38	8.2	0 . 8
7.	15.42722	80.39472	1735	867	35	10.2	-1.2
₽ •	15.37000	80.17806	1853	482	37	8.8	0.2
G a	15.16139	80.08833	2144	400	37	8.8	0.2
/)	15,10222	80.05333	2229	358	38	8,2	0 * 8
**	16,23333	81,68999	435	2795	35	10,2	-1.2
2.	15.90899	81.20888	909	2171	38	8.2	0.8
3.	16.26567	81.46567	418	2530	36	9.5	-0.5
1.	15,65083	80,77972	1365	1499	37	8.8	0.2
5.	15.39528	80.07611	1840	279	37	8.8	0.2

For the water depth of 9 m

r = -0.99902

Se = 0.684 m

 $\sigma = 0.706 \text{ m}$ 

For the water depth of 18 m

r = -0.995

Se = 1.255 m

 $\sigma = 1.2828 \text{ m}$ 

The negative sign shows the reciprocal relation. Here we find that standard error for 18 m water depth is more than that of 9 m water depth. It may be due to the reason that for same difference in reflectance values of deep and shallow water, the variation in water depth is more for deep water than for shallow water.

From the correlation analysis, it is clear that coefficient of correlation is about one, which shows that results obtained are quite satisfactory.

## 4.4 DRAWING OF WATER DEPTH CONTOUR

There are two methods for drawing contours

- (a) Direct Method
- (b) Indirect Method

In direct method of contouring, contour to be plotted is actually traced on the ground. We are concerned about only those points which lie on the contour. In the indirect method, some suitable guide points are selected and guide points need not necessarily be on the contours. The guide points serve as a basis for the interpolation of contours. In linear interpolation, it is assumed that water depth variation between two guide points is proportional to distance.

In the present work, indirect method of contouring was done. Total area was divided into small grids by drawing lines parallel to the longitudes and latitudes. The interval between the lines drawn parallel to latitude and longitude are kept 4 min and 2 min respectively. The latitudes and longitudes of all the corner points of the grids were calculated. The corresponding line numbers and pixel numbers were calculated by previously mentioned program. The reflectance values of all the pixels corresponding to grid corners were calculated and finally the depth of those points were calculated. The values are tabulated in Table 4.3. Points of equal depths (9 m and 18 m) were calculated by interpolation and contours were drawn. The contours what we obtained, is more or less same as the contours of the toposheet.

TABLE: 4.3

The season sees now sees need to see the seeson need to seeson need to see the seeson need	TO A T I T U D E	LONGITUDE	LTNE No.	PIXEL No.	REFLECTANCE VALUE	WATER DEPTH
1234567E901234567890123456789012345678901234 1111111117222222222333333333344444	3007307307307307307307307307307307307307	811.8811.887777777777777777777777777777	2604715047158237158263715826937171482693726153 233344533444555334455566773344455566773	72605937159482160493726159382615715048271592 1111112000000189999997777888888866666777775 13333333333333333222222222222222222222	58967546360754313975498962988654238542977542 543322543322254322222254333222222225433322225	58753735757373368437378757411273672297433736 13790781594078261607803792688378258002600782 223 1223 11223

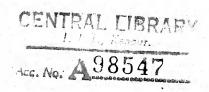


TABLE: 4.3 (....CONTINUED)

1 4 C m	Harman and the same of the sam	LONGITUDE	LINE No.	Plxet.	REFLECTANCE VALUE	WATER DEPTH
0078901234567890123456789012345677777777777777777888888888999999999999	######################################	77777777777777777777777777777777777777	48269370482594825937049159315936049159260479260471582 7160493626154837160593826104837260593827155049372605 3445556677883344555666778889444556677788899004555667788	6049371604837159482605937260482715948260595937160482555555666666334444444555555522232222222222	37 441098865419375532098766414847652199876541963930764	7899874112733478287741322330838527844132733753471859 -************************************

TABLE: 4.3

1 to	DATITOE.	LONGITUDE	LINE No.	PIXEL No.	REFLECTANCE VALUE	WATER DEPTH
	3307307307307307307307307307307307307307	77777777777777777777777777777777777777	6037158237158269371482693937048259360048159360471260 9482715004937260594827150726159482716826159483716150 89900111266677888999001112233899001112233344223 111111 1111111 111111111111111111111	7159372626049371594826049260482615937493715937260048 912355689780124578912457893467801134578346790134679890 23333333330011111111222222222222222222	20986654952843186310987649183762987544826296296228654606 33222225554444333333222225544333222222554453443	7741227377683762588741323798785741373065275771273515

TABLE: 4.3 (....CONTINUED)

The state of the s	and the same of th	LONGITUDE	LINE No.	PIXEL	REFLECTANCE VALUE	WATER DEPTH
TILLIA   T	150-6007300730073007300730073007300730073007	00003333333337777777000000333333333333	11111111111111111111111111111111111111	26082604937159377169371159371593726048604826048 2999777688666667755555693333333444444455122222223333333333333333	22244316548637754848275482831976208765481849865419876 222443322244432224443322255444333333322222255444332222	1733782738373832737300876785771327309837252984132 123 1223 1223 1223 1223

TWATE: : 4.3 (....C.JALIAUED)

101123456789012345678901234568 1000000000111111111120222022				E . 12600482604826048260489057891245689124568912456891245689124568912456891245689124568912456891245689124568912456891245689128888	I A SELECTION OF THE PROPERTY	WE 1222 11111222223
3301204567890423466789042M456789042A4444555	3073073073073073073073073073073073073073	00000000000000000000000000000000000000	771482594837452746537825484594837465374837444455557	93604826048260715937159371 00000011111111128888999999999 1111111111111	TOTALOGO BITTIONIO MOMBIO MENOMENTO PORTUGO DE LA COMPANSIONIO DELIGIO DELIGIO DE LA COMPANSIONIO DE LA COMPANSIONIO DELIGIO DE LA COMPANSIONIO DE LA COMPANSIONI DELIGIO DE	7 9 9 0 1 24 8 0 3 3 7 7 8 1 2 1 4 5 7 8 0 2 2 3 4 5 7 8 0 2 2 3 4 5 7 8 0 2 2 3 3 4 5 7 8 0 2 2 2 2 2 2 3 4 5 7 8 0 2 2 2 2 2 2 3 4 5 7 8 0 2 2 2 2 2 3 4 5 7 8 0 2 2 2 2 2 2 3 4 5 7 8 0 2 2 2 2 2 2 3 4 5 7 8 0 2 2 2 2 2 2 2 3 4 5 7 8 0 2 2 2 2 2 2 3 4 5 7 8 0 2 2 2 2 2 2 3 4 5 7 8 0 2 2 2 2 2 2 2 3 4 5 7 8 0 2 2 2 2 2 2 2 3 4 5 7 8 0 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2 2

TABLE: 4.3 (....CONTINUED)

					•	
SI.	LATITUDE	LONGITUDE	LINE No.	PIXEL No.	REFLECTANCE VALUE	WATER DEPTH
	0730730730730730730730730730730730730730	77777000000000000000000000000000000000	11111121111111111111111111111111111111	1245789346790134678013457803567902346790134678056890 000000777777888888889999999906666777777778888888899555556	9876554978864297641998766547488429742119876655472630 22222255444433333222222222554444333333322222222	4132773728853278598441322732088327884132277326571 ••••••••••••••••••••••••••••••••••••

TABLE: 4.3 (....CONTINUED)

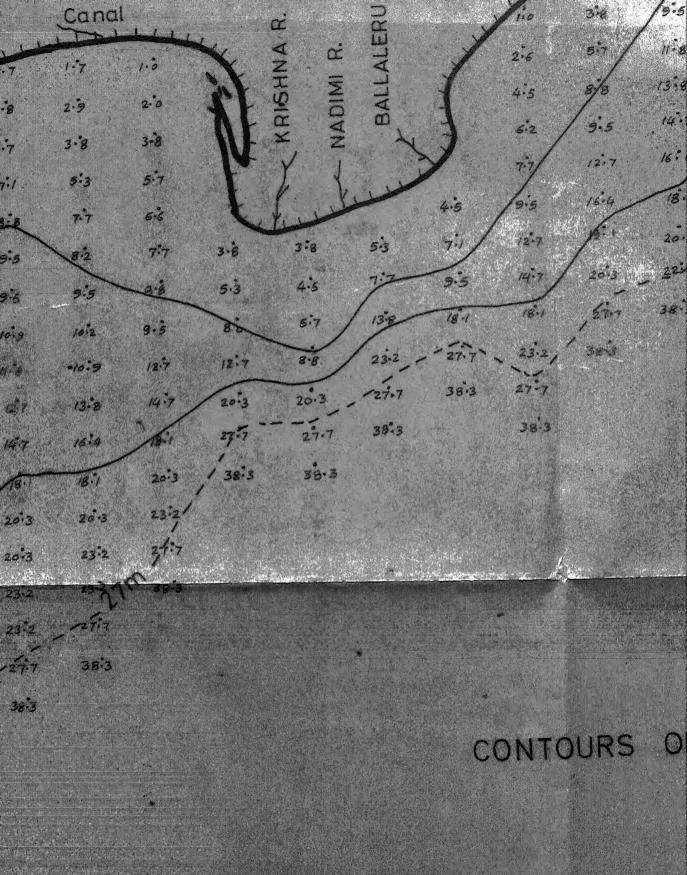
SI. No.	LATITUDE	LONGITUDE	LINE No.	PIXEL No.	REFLECTANCE VALUE	WATER
56789042345678901234567890123456789012345678901234567890123456789012345678901234567890123533333333333333333333333333333333333	3063063073073073073073073073073073073073073073	10000000000000000000000000000000000000	26047159260371593704825937046047159260475937 111111122217788899900112223899900112233 1111112222378889900112223899900112233 122337260574827112223899990011223312223	37159370482604715937158260489371592604820482 23567902346790912356890235679124568912357891 266666677777778455555555566666344444444555553334	87631098776655360755210099879063109876547413	28588741332277351822787744137257617285292093 8888134680033772478002344466800345677889001222 111111222222

16°45

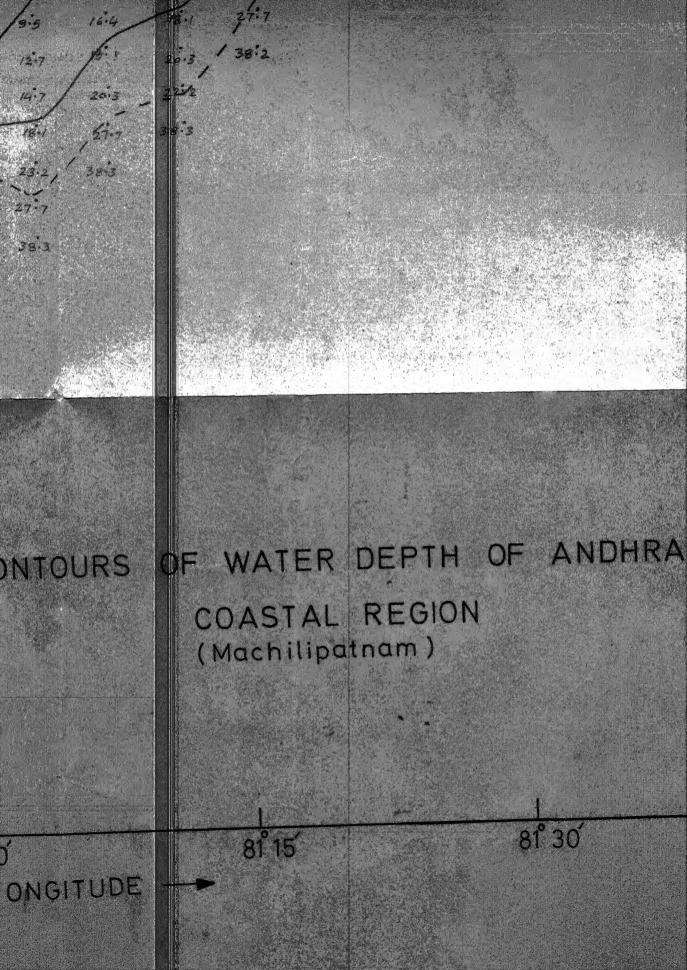
1630

16 15 -

S.GODAWARI R. 0.7 2.6 2.6 . 3-8 2.3 3.8 5·3 7:1 5.7 6.6 7.7 8:2 9.5 11.8 9:5 10:2 10:9 20.3 16.4 14.7 12-7 10:9 10.9 127.7 20.3 16.4 20.3 12.7 13:8 38.3 121.7 27.7 18.1 عزوا 14.7 7 38.3 38-3 18:1 20.3 16:4 7 23.2 20.3 18.1 127.7 27.7 18:1 38.3 38.3 23.2 .3 ,27.7 . 3.2 38.3 5.2 / 38·3



				9		WIN
			سلا	0.3	2.6	2.6
			0.3	<b>3.4</b>	5.7	5.3
		6.3	20	6.7	_BiB	8:2
	1	0.7	3:8	18.8	10:9	10.2
	1	1:5	5:3/	10.2	16.9	10:9
MACHILIPATNAM	0:7	2:3	8/8	11.8	13:8	12.7
MACHIELLA	7:4	3:4	9.5	12.7	14.7	16:4
I	2:6	5:7	10.2	14.7	16:4	Zö.3
4	3 <b>:</b> g	7.1	12.7	16.4	/18·1	20.3
KRISHNA R.	<b>5</b> :3	8:/8	13.8	1811	18-1	27.7
	5.7	9:5	16.4	20.3	23.2,	38:3
10.7	6.6	10.9	16.4	25-2	,27.7	
2.9	2.7	12.7	y h	23.2	CONTRACTOR OF THE PROPERTY OF	
Z 26	19:5	14.7	120.3	38.3		
2.6 57/	11:8	16:4/	23.4			
1 1 45 8/8	13:8	/8-1	z4.7			
W 1 6:2 / 9:5	14.17	23:2	1 38.3			
7 77 12.7	16:4/	23:2/				
4:5 /9:5 16:4	18:1	34:7				
53 41 / 124	20.3	/ 38:2				
	2212					
	38-3					
13.8 18.1 18.1						



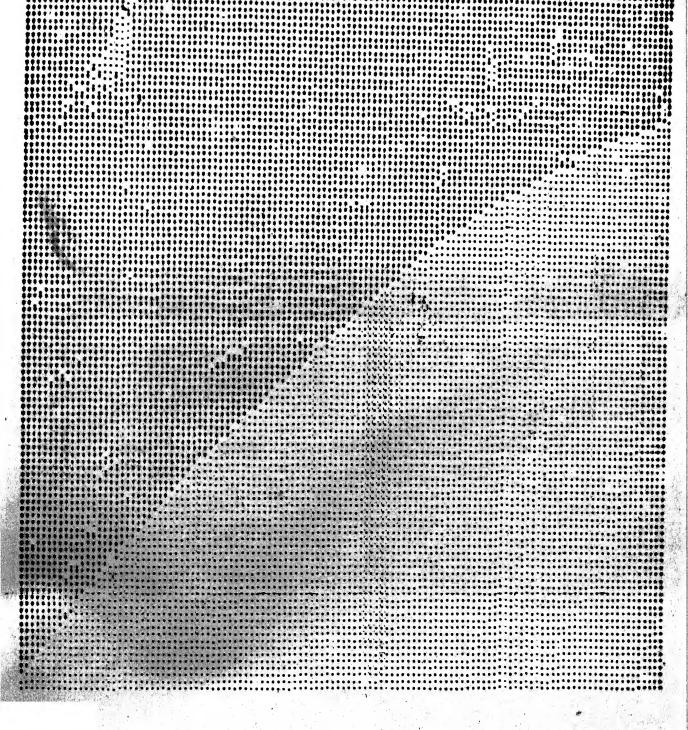


Fig. 4.2: Map of Coastal Area

\*→Land ...→ Water

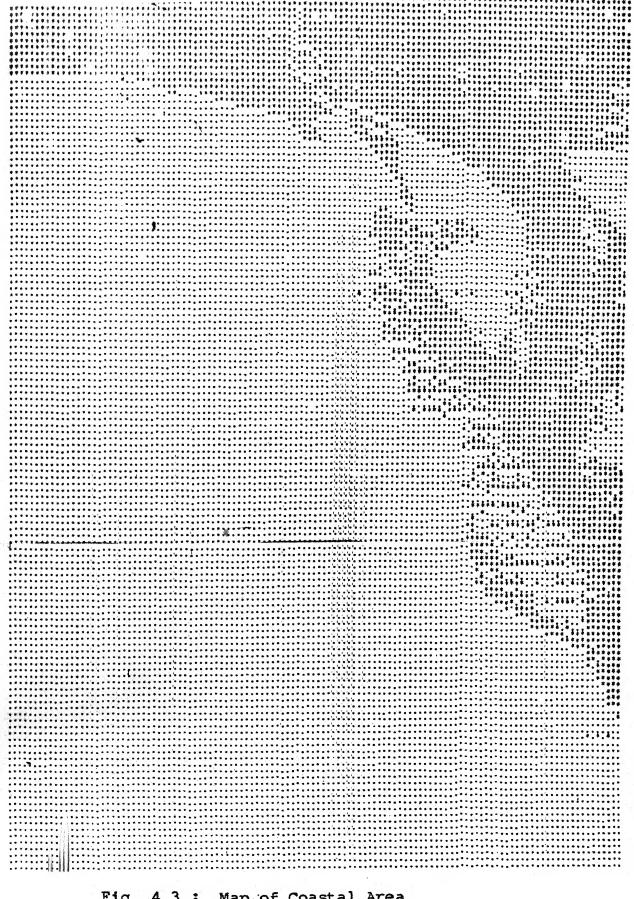


Fig. 4.3 : Map of Coastal Area

\* → Land ... → water

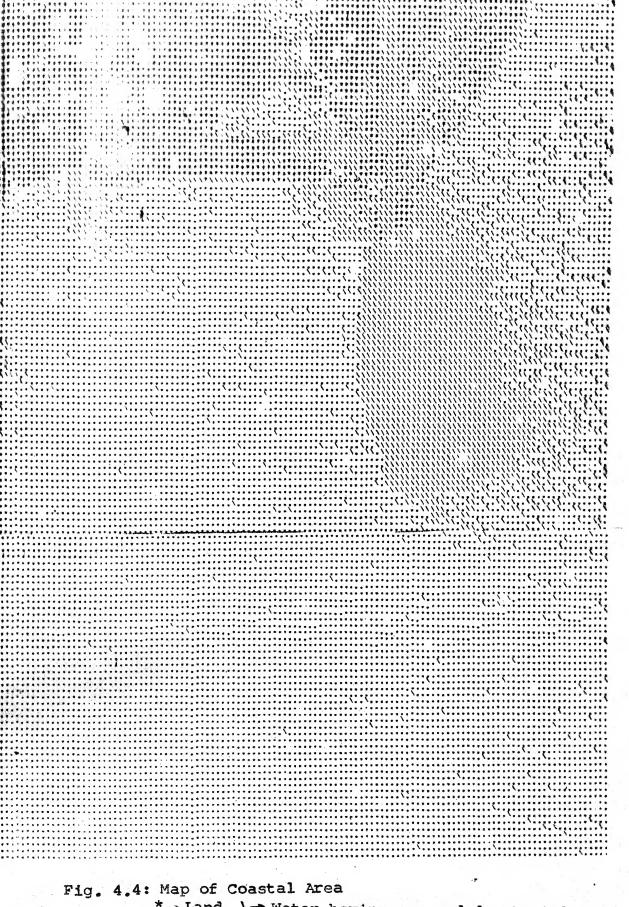


Fig. 4.4: Map of Coastal Area

\*→Land \→ Water having suspended materials
..→Clear Water

#### CHAPTER 5

# CONCLUSIONS, LIMITATIONS AND RECOMMENDATIONS FOR FUTURE WORK

#### 5.1 CONCLUSIONS

The following conclusions are made after studying the Landsat data of the study area.

Bathymetry and coastal geomorphology can now be obtained from satellite data using Remote Sensing techniques specially in Band 1 (Landsat 4 and 5) where water is shallow and clear. The information obtained by this Remote Sensing technique is much time saving and cost effective compared to other conventional method like sounding machine, echo sounding, sonar, visible spectral range aerial photographic techniques. The scope of this technique becomes more, particularly in those area where no recent hydrographic surveys are available. This project introduce the newly developed technique of 'Bathymetry' i.e. delineation of coastal feature, water depth, water having suspended materials, et with the help of spectral reflectance by sensor.

The scope of present project enables us to obtain the preplanned position and to improve the internal relative position of the coastal features, in detecting features missed by earlier

surveyors, in improving delineation of features and to demarcate high water line and low water line.

Remote Sensing technique of Bathymetry is highly useful in areas of mobile seabed which are changing frequently like sandspits, at mouth of rivers, newly formed banks, islands to determined the rate of change of features and to update the information. It further helps in the study of coastal erosion, bank, channel, ground, submerged feature etc. The delineation of features by remote sensing is not only time saving and cost effective but also helps in developing new ports and sea routes.

In conclusion, it may be stated that present work helps in the field of hydrography as follows:

- (a) To locate the high water line and low water line.
- (b) To get water depth with good accuracy.
- (c) To detect the character of coast and foreshore.
- (d) To detect uncharted, mispositioned and submerged features which are hazardous to navigators.
- (e) To define base line for territorial boundaries.
- (f) To locate newly formed islands, banks, shoals etc.
- (g) To show frequent change in ports, harbours, coastal areas of which immediate survey is not possible.

## 5.2 LIMITATIONS

The followings are requirements of this study:

- 1. The area should be cloud free.
- 2. Water should not have suspended materials because it can alter the reflectance value.
- It should be shallow water.
- 4. There should not be change in refractive index of sea water. For our work, it is assumed that refrective index of sea water is same all over the sea.
- 5. The alignment of canals, drains, streams can be well interpreted but depth of water in canal, streams can not be found from Landsat Data.
- 6. Landsat data is found useful for regional study of large water bodies rather than small areas.
- 7. In general, ponds larger than 1 hectare and streams wider than 20 m are seen easily by Landsat imageries but individual objects less than 70 m can not be distinguished easily by Landsat data.

# 5.3 RECOMMENDATION FOR FUTURE WORK

- The work can be further extended so that water depth can be predicted for deeper water.
- If water depth is calculated keeping in mind that refrective

index of sea water is changing from place to place, we can get better result.

In this project, contouring was done for clear water.
 One may extend this for water having suspended materials.

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# CC APPENDIX : 1 000000 FILE NAME: MAN.FOR PROGRAM TO CONNVERT THE GEOGRAPHICAL COORDINATE TO ITS CONICAL ORTHOMORPHIC COORDINA X & Y THE CONSTANTS ARE VALID FOR STUDY AREA ONLY. IMPLICIT DOUBLE PRECISION(A-H,K-Z) DIMENSION LAT (150),LONG(150),LATRAD(150) LANGD=79.8428899 LATD=14.753744122\*3.1415921/180.0 A=6377277.6 ELLIP=10/300.8 ELLIP=10/300.8 ELLIP=ELLIP\*ELLIP SIN=DSIN(LATO) SINSQ=SIN\*SIN DENDM=(1.0-(ELLIP\*SINSQ)) NO=A/(DSQRT(DENOM)) RD=NO\*((1.0-ELLIP)/DENOM) P=NO\*((1.0-ELLIP)/DENOM) READ(21,\*)II DD 100 1=1,II READ(21,\*)DEG1,MIN1,SEC1,DEG2,MIN2,SEC2 LAT(I)=DEG2+MIN2/60.0+SEC1/3600.0 LONG(I)=DEG1+MIN1/60.0+SEC1/3600.0 LATRAD(I)=LAT(I)\*3.1415926/180.0 LONGDF=LONG(I)-LANGO GAMMA=LONGDF\*SIN GAMMA=GAMMA\*3.1415926/180.0 M1=M\*\*3/(6.0\*RC\*NO) TAN=DSIN(LATO)/DCOS(LATO) M1=M\*\*43/(6.0\*RC\*NO) TAN=DSIN(LATO)/DCOS(LATO) M1=M\*\*43/(6.0\*RC\*NO) M1+M\*\*43/(6.0\*RC\*NO) M2-M1\*\*(M\*\*TAN)/(4.0\*NO)) M1-M\*\*43/(6.0\*RC\*NO) M2-M1\*\*(M\*\*TAN)/(4.0\*NO)) M1-M\*\*43/(6.0\*RC\*NO) M2-M1\*\*(M\*\*TAN)/(4.0\*NO)) M1-M\*\*43/(6.0\*RC\*NO) M1-M\*\*43/(6.0\*RC\*NO) M2-M1\*\*(M\*\*TAN)/(4.0\*NO)) M1-M\*\*43/(6.0\*RC\*NO) M1-M\*\*43/(6.0\*RC\*NO) M1-M\*\*43/(6.0\*RC\*NO) M1-M\*\*43/(6.0\*RC\*NO) M1-M\*\*43/(6.0\*RC\*NO) M1-M\*\*43/(6.0\*RC\*NO) M1-M\*\*43/(6.0\*RC\*NO) M

STOP

# APPENDIX : 2

```
CCCCCCCC
                 FILE NAME : L
THIS PROGRAM
POINT ON THE
NUMBER ON THE
COORDINATE NTS
THE CONSTANTS
                                        LINPIX FOR CONVERTS THE LATITUE EARTH'S SURFACE TO E IMAGERY AFTER CONV
                                                                              DE
                                                                                    AND LON
                                                                                                 GITUDE O
UMBER AN
CONICAL
                                                                                                                 F A GIVE
D PIXEL
ORTHOMOR
                                                                       CONVERSION
             TS ARE VALID FOR THE STUDY
1001
100
10
                                   LATITUDE=',F23.16.5X, LONGITUDES=',F23,9(1H=)//3X, LATITUDE',5X, LONGITUDE',)
20
```

```
CC
                            APPENDIX : 3
 CCCCCC
        OF THE OF CCT DEC+10 *****
                                           CCT
THIS IS VA
COMPATIBLE
*******
                                                     VALID
      , MODE="DUMP", RECORD SIZE=1125, DET
11
1200
                                  , MODE = DUMP : , RECORD SIZE=1125
21
101
10
```

```
PIMENSION VX(10),X(10),AT(3,10),A(10,3),AL(10),R1(10),R2(10)
DIMENSION X1(10)
OPEN(UNIT=22,DEVICE='DSK',FILE='FOR22.DAT',ACCESS='SEGOU')
OPEN(UNIT=21;DEVICE='DSK',FILE='FOR21.DAT',ACCESS='SEGIN')
Y=-2,267
                OPEN CUNIT=21
Y=-2.267
V0=55
V0=17
CK=0.075
DO 56 K=1,12
VD=VD
CK=CK
V0=VO
DO 10 I=1,10
ICK=K
                Vomvo
DO 10 I=1,10
ICK=K
IF(ICK.EQ.1)READ(21,*)X(I),VX(I)
READ(21,*) X(I),VX(I)
R1(I)=-ALOG((VX(I)-VD)/(V0-VD))/(Y*CK**2)
R2(I)=(VX(I)-V0)/(Y*CK*(V0-VD)*(VX(I)-V0))
R3(I)=-(VX(I)-VD)/(Y*CK*(V0-VD)*(VX(I)-V0))
AL(I)=X(I)
X1(I)=ALOG((VX(I)-VD)/(V0-VD))/Y*CK
AL(I)=X1(I)-X(I)
TYPE*,VX(I),X(I),R1(I),R2(I)
TYPE*,AL(I)
CONTINUE
I=1
                I=1
DQ 30 J=1,10
AT(I,J)=R1(J)
I=2
DQ 40 J=1,10
               DO 40 J=1,10
AT(I,J)=R2(J)
I=3
DO 45
                DO 45 J=1,10
AT(I,J)=R3(J)
J=1
                DO 31 I=1,10
A(I,J)=R1(I)
J=2
DO 41 I=1,10
A(I,J)=R2(I)
                           41, I=1(10
,J)=R2(1)
                DO 41 I=1,10

A(I,J)=R2(I)

J=3

A(I,J)=R3(I)

TYPE*,(AT(I,J),I=1,3),J=1,10)

TYPE*,(AT(I,J),I=1,10),J=1,3)

WRITE(22,*) (VX(I),X(I),R1(I),R2(I),I=1,10)

WRITE(22,*) ((AT(I,J),I=1,3),J=1,10)

WRITE(22,*) ((A(I,J),I=1,3),J=1,3)

L=3

M=10

N=3

CALL MULT(AT,A,B,L,M,N)
                 CALL MULT(AT, A, B, L, M, N)
                 N=3
                 CALL INVERT(B, N)
```

```
L=3
                  M = 3
                  N=10
               CALL MULT(B, AT, C, L, M, N)
                  L=3
               M=10
N=10
CALL MULT(C, AL, D, L, M, N)
D1==D1
D2==D2
D3==D3
CK=CK+D1
VD=VD+D2
V0=V0+D3
WRITE(22,*) I. CK, VD, V0, D1, D2, D3
FORMATIC(BX, 'ITERATION NO=','21/,8X,'CK=',F10.5/,8X,'VD=',F10.5/,8X,'D2=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'D3=',F10.5/,8X,'
               1 = 1
DO 55 J=1, L

SUM=0.0

DO 46 K=1, M

SUM=SUM+R(I, K)*S(K, J)

T(I, J)=SUM

RETURN

END

SUBROUTINE FOR INVERSE OF MA'S

SUBROUTINE INVERSE OF MA'S

SUBROUTINE INVERT(A, N) EX(50)

DO 113 I=1, N

INDEX(I)=0

AMAX=-1:0

DO 115 I=1, N

IF(IP=ABS(A(I, I)) 1.15, 1.15, 1.15

TEMPEABS(A(I, I)) 1.15, 1.15, 1.17

IF(IP=ABS(A(I, I)) 1.15, 1.15, 1.17

IF(IP=ABS(A(I, I)) 1.15, 1.15, 1.17

ICOL=I

AMAX=TEMP

CONTINUE I, ICOL)

IF(ICOL)=1, ICOL)

A(ICOL, J)=1, N

A(ICOL, J)=
                     RETURN
                     END
```

END

```
CC
                                                                                                                       APPENDIX : 6
  00000000
              *****************
             INTEGER IPIX(960), INTMAT(960), GREY(10)
INTEGER CONTOR(10), INTVL(10), UPLIM(10), UPLIM(10)
OPEN(UNIT=25, DEVICE='DSK', FILE='INPUT')
OPEN(UNIT=6, DEVICE='DSK', FILE='MAP1')
OPEN(UNIT=7, DEVICE='DSK', FILE='MAP2')
OPEN(UNIT=8, DEVICE='DSK', FILE='MAP4')
OPEN(UNIT=9, DEVICE='DSK', FILE='MAP4')
OPEN(UNIT=10, DEVICE='DSK', FILE='MAP4')
OPEN(UNIT=11, DEVICE='DSK', FILE='MAP6')
OPEN(UNIT=12, DEVICE='DSK', FILE='MAP6')
OPEN(UNIT=13, DEVICE='DSK', FILE='MAP8')
                               INTERACTION BEGINS
                              TYPE 10
FORMAT(' ALL TERMINAL INPUT IS FORMAT FREE',',',
18X, TYPE IN THE NUMBER OF CLASSES FOR SLICING')',
ACCEPT *, NCLASS
TYPE 20
FORMAT(8X, TYPE IN CHARACTERS FOR REPRESENTATION
1 OF THE CLASSES')
ACCEPT 25, (GREY(I), I=1, NCLASS)
FORMAT(10Å1)
   20
20000
                              TO GET THE CONTOUR LEVEL AND FROM TERMINAL
                                                                                                                                  INTERVAL OF
                             DO 30 I=1, NCLASS

TYPE 40.I

FORMAT(8x.'TYPE IN THE CONTOUR LEVEL AND INTERVAL', 12)

ACCEPT *. CONTOR(I), INTVL(I)

UPLIM(I)=CONTOR(I)+INTVL(I)

LOWLIN(I)=CONTOR(I)-INTVL(I)

CONTINUE

TYPE 50

FORMAT(8x,'TYPE IN THE LENGTH AND BREADTH OF INPUT')

ACCEPT *. NLINE, NPIX

TYPE 55

FORMAT(8x,'TYPE IN THE NUMBER OF OUPUT FILES')

ACCEPT *, NOUTFL
   40
   30
```

INTERACTION ENDS.

```
CC
CC
CD O INITIALISE THE INTERMEDIATE MATRIX INTHAT

DO 65 I=1,NPIX
INTMAT(I)='
CC
TO READ AND CLASSIFY THE INPUT DATA

CILOOP=0
CONTINUE
READ(25,*)(IPIX(I),I=1,NPIX)
ILOOP=ILOOP + 1
DO 70 J=1,NPIX
DO 70 J=1,NCLASS
IF((IPIX(I),LE.UPLIM(J)).AND.(IPIX(I).GE.LOWLIM(J)D)
INTMAT(I)=GREY(J)
CONTINUE
CC
To transfer the results from INTMAT to output files
for getting the line printer map.

NUNIT=6
M=1
N=M+119
DO 80 I=1,NOUTFL
WRITE(NUNIT,90)(INTMAT(J),J=M,N)
FORMAT(120A1)
NUNIT=NUNIT+1
M=N+1
M=
```